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## Atmospheric Research

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## Radiation balance in a cloudy atmosphere with account for the 3D effects

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## ABSTRACT

In this paper we derive radiation balance equations for the cloudy atmosphere with account for horizontal inhomogeneity effects. It is assumed that the clear sky and cloud occupy two different half-spaces separated by a vertical plane. Numerical results for corresponding fluxes are derived using the 3D code RADUGA based upon the discrete ordinate method. It is shown how "jump" of optical properties forms shadowing or brightening at both the cloud edge and the neighboring clear sky.

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## 1. Introduction

The spatial distribution of brightness of broken cloud fields is very inhomogeneous both for reflected and transmitted solar light (Marshak et al., 2000; Varnai and Marshak, 2002). Clearly, this is the case for the absorptance as well. A humid aerosol layer surrounding the cloud could be a factor (Redemann et al., 2009). However, the careful analysis of the problem indicates that the 3D radiative interaction between clouds and surrounding clear areas is a major factor explaining corresponding extrema (Varnai and Marshak, 2009). Studies of solar light propagation based on the discrete ordinate method (Bass et al., 1986) (code RADUGA) in 2D regions, containing both cloud and clear sky (aerosol) with one vertical common boundary, shows (Nikolaeva et al., 2005) that brightness extrema in the vicinity of a cloud edge occur even if an intermediate aerosol layer is absent.

The origin of extrema of the intensity of solar light reflected and transmitted by the cloudy atmosphere is studied in this paper. Only one factor, namely, the horizontal

variations of optical properties of the atmosphere is taken into account. To eliminate the influence of intermediate layers, we consider a model problem, where a cloud and an aerosol (clear sky) have only one common (vertical) boundary  $x = 0$  (see Fig. 1). Besides, both media are assumed to be homogeneous. Hence, the 3D radiative effects occur only in the vicinity of the cloud edge. Also it is clear that far from the common boundary both reflected and transmitted light intensities depend only on the height  $z$ .

We simulate the radiative transfer in atmosphere using the 3-D transport equation and the discrete ordinate method (Section 2). Applying the radiative balance relation for the whole spatial region, not for a single pixel, see Section 3, it is shown how the horizontal radiative transfer caused by a "jump" of optical properties of a medium under consideration forms shadowing or brightening at both cloud edge and neighboring clear sky (Section 4).

## 2. Mathematical model

Let us put the coordinate origin at the top of the atmosphere and direct the axis  $z$  to the ground. We place the axis  $y$  along the boundary of the cloud and aerosol (in the horizontal direction, see Fig. 1). It is assumed that the illumination direction is perpendicular to the axis  $y$  (the azimuth  $\phi$  of the

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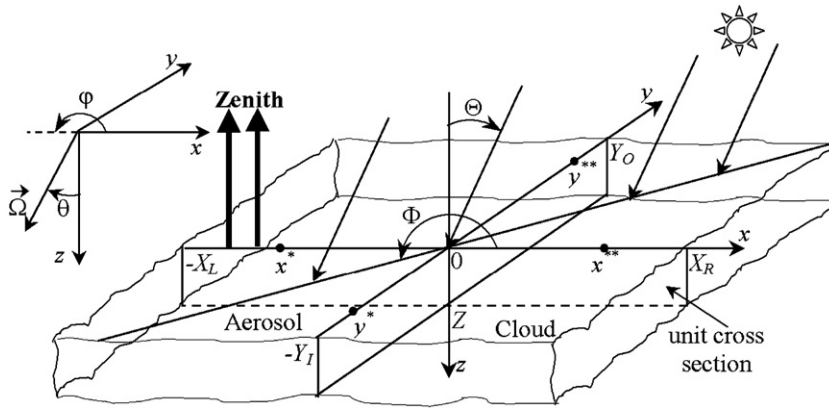


Fig. 1. The studied atmospheric model.

sun is equal to  $0^\circ$  or  $180^\circ$ , see Fig. 1). In this case the radiation intensity does not depend on coordinate  $y$  and can be found as a solution of the 2D radiation transport equation in a cross section of the layer  $-X_L < x < X_R$ ,  $0 < z < Z$ , having unit thickness over coordinate  $y$ :

$$\xi \frac{\partial I(x, z, \theta, \varphi)}{\partial x} + \beta \frac{\partial I(x, z, \theta, \varphi)}{\partial z} + \sigma_{\text{ext}}(x, z) I(x, z, \theta, \varphi) = \frac{1}{4\pi} \sigma_{\text{ext}}(x, z) \omega_0(x, z) \int_0^\pi \sin \theta' d\theta' \int_0^{2\pi} I(x, z, \theta', \varphi') p(x, z, \chi(\theta, \varphi, \theta', \varphi')) d\varphi', \quad (1)$$

where

$$\xi = \sin \theta \cos \varphi, \quad \beta = \cos \theta, \quad \chi(\theta, \varphi, \theta', \varphi') = \cos \theta \cos \theta' + \sin \theta \sin \theta' \cos(\varphi - \varphi'), \quad -X_L < x < X_R, \quad 0 < z < Z, \quad 0 < \theta < \pi, \quad 0 < \varphi < 2\pi. \quad (2)$$

The function  $I(x, z, \theta, \varphi)$  in Eq. (1) defines the light intensity at the spatial point  $(x, z)$  in the direction  $\Omega$ , determined by the angles  $(\theta, \varphi)$  (see Fig. 1). Functions  $\sigma_{\text{ext}}(x, z)$  and  $\omega_0(x, z)$  are extinction coefficient and single scattering albedo, respectively. The scattering phase function  $p(x, z, \chi)$  is normalized by the following relation

$$\frac{1}{4\pi} \int_0^\pi d\theta' \sin \theta' \int_0^{2\pi} d\varphi' p(x, z, \theta, \varphi, \theta', \varphi') = 1. \quad (3)$$

The following boundary conditions for Eq. (1) are used:

$$I(x, 0, \theta, \varphi) = F_0 \delta(\cos \theta - \cos \Theta) \delta(\varphi - \Phi) \text{ as } \cos \theta > 0, \quad (4)$$

$$I(x, Z, \theta, \varphi) = 0 \text{ as } \cos \theta < 0, \quad (5)$$

$$I(X_R, z, \theta, \varphi) = I(x^*, z, \theta, \varphi) \text{ as } \cos \varphi < 0, \quad (6a)$$

$$I(-X_L, z, \theta, \varphi) = I(x^{**}, z, \theta, \varphi) \text{ as } \cos \varphi > 0. \quad (6b)$$

Eq. (4) means that the medium under consideration is illuminated by the wide solar light beam at the top (the zenith angle  $\Theta$  and azimuth  $\Phi$ , see Fig. 1). The irradiance at a unit area perpendicular to the solar beam is assumed to be

equal to  $F_0$ . Eq. (5) means that the bottom boundary  $z = Z$  is black.

The coordinates  $x^*$  and  $x^{**}$  in periodic conditions (6a) and (6b) correspond to points, put far off both the exterior boundaries  $x = -X_L$ ,  $x = X_R$  and the interior boundary  $x = 0$ , see Fig. 1. The conditions (6a) and (6b) and also the homogeneity and semi-infinity of each medium guarantee independence of solution  $I(x, z, \theta, \varphi)$  on coordinate  $x$  far away from the interior boundary. The radiation intensity far away from the interior boundary can be obtained via the 1D slab layer model.

Let  $I^{\text{aer}}(z, \theta, \varphi)$  and  $I^{\text{cloud}}(z, \theta, \varphi)$  be the light intensity in clear sky and cloud, accordingly. Note, that the radiation intensity in the framework of IPA (Independent Pixel Approximation) model not taking into account the radiative exchange between media is defined as

$$I^{\text{IPA}}(x, z, \theta, \varphi) = \begin{cases} I^{\text{aer}}(z, \theta, \varphi) & \text{as } x < 0, \\ I^{\text{cloud}}(z, \theta, \varphi) & \text{as } x > 0. \end{cases} \quad (7)$$

A boundary layer is defined in this paper as a spatial region neighboring the boundary of both media, where the error of the IPA solution is large. The sizes of boundary layers  $x^{\text{aer}}$  and  $x^{\text{cloud}}$  are defined via the following expressions

$$\max_{x < -x^{\text{aer}}} \max_{z, \theta, \varphi} |\varepsilon(x, z, \theta, \varphi)| = \varepsilon_0 \text{ and } \max_{x > x^{\text{cloud}}} \max_{z, \theta, \varphi} |\varepsilon(x, z, \theta, \varphi)| = \varepsilon_0. \quad (8)$$

Here

$$\varepsilon(x, z, \theta, \varphi) = 1 - I(x, z, \theta, \varphi) / I^{\text{IPA}}(x, z, \theta, \varphi), \quad (9)$$

$\varepsilon_0$  is the relative error. For example,  $\varepsilon_0$  can be chosen as an instrumental error. In a similar way sizes of boundary layers can be estimated for both light intensity  $I(x, z, \theta, \varphi)$  and some other radiative characteristics (for instance, reflectance and transmittance, see Section 3).

### 3. Radiation balance

We derive the balance equation for the whole region, rather than just for one pixel (Titov, 1998; Widłowski et al., 2006). Let introduce functions, defining radiation fluxes at the vertical boundaries of the section  $[-X_L, X_R] \times [0, Z]$  of unit width:

$$j_R(z) = \int_0^\pi d\theta \sin\theta \int_0^{2\pi} d\varphi \sin\theta \cos\varphi I(X_R, z, \theta, \varphi) \quad (10a)$$

—at the right boundary  $x = X_R$ ,

$$j_L(z) = - \int_0^\pi d\theta \sin\theta \int_0^{2\pi} d\varphi \sin\theta \cos\varphi I(-X_L, z, \theta, \varphi) \quad (10b)$$

—at the left boundary  $x = -X_L$ .

Each of these functions is positive if the main radiation flux across corresponding spatial point of a boundary is directed out of the section and negative in the opposite case.

We introduce fluxes of radiation leaving the section via its top and bottom boundaries

$$j_T(x) = \int_{\cos\theta < 0} d\theta \sin\theta |\cos\theta| \int_0^{2\pi} d\varphi I(x, 0, \theta, \varphi) \quad (11a)$$

—for the top boundary,

$$j_B(x) = \int_{\cos\theta > 0} d\theta \sin\theta \cos\theta \int_0^{2\pi} d\varphi I(x, Z, \theta, \varphi) \quad (11b)$$

—for the bottom boundary.

The absorptivity  $c$  at a given point  $M(x, z)$  can be found using the following equation:

$$c(x, z) = \sigma_{abs}(x, z) \Phi(x, z), \quad (12)$$

where  $\sigma_{abs} = (1 - \omega_0) \sigma_{ext}$  is the absorption coefficient and

$$\Phi(x, z) = \int_0^\pi d\varphi \int_0^\pi d\theta \sin\theta I(x, z, \theta, \varphi) \quad (13)$$

is the actinic flux.

To obtain the radiative balance relation we integrate the Eq. (1) over angles  $\theta$  and  $\varphi$  and the spatial variables  $x, z$  in the section  $[-X_L, X_R] \times [0, Z]$  of unit width, see Fig. 1, taking into account the normalization equality (3) and boundary conditions (4)–(6b). In particular, the integration of the first term on the left side leads to the expression

$$\begin{aligned} & \int_0^Z dz \int_{-X_L}^{X_R} dx \int_0^\pi d\theta \sin\theta \int_0^{2\pi} d\varphi \sin\theta \cos\varphi \frac{\partial I}{\partial x} \\ &= \int_0^Z dz \int_0^\pi d\theta \sin\theta \int_0^{2\pi} d\varphi \sin\theta \cos\varphi [I(X_R, z, \theta, \varphi) - I(-X_L, z, \theta, \varphi)] \\ &= \int_0^Z dz \{j_R(z) + j_L(z)\} \end{aligned} \quad (14)$$

It follows for the second term

$$\begin{aligned} & \int_0^Z dz \int_{-X_L}^{X_R} dx \int_0^\pi d\theta \sin\theta \int_0^{2\pi} d\varphi \cos\theta \frac{\partial I}{\partial z} \\ &= \int_{-X_L}^{X_R} dx \int_0^\pi d\theta \sin\theta \int_0^{2\pi} d\varphi \cos\theta [I(x, Z, \theta, \varphi) - I(x, 0, \theta, \varphi)] \\ &= \int_{-X_L}^{X_R} dx \left[ - \int_{\cos\theta > 0} d\theta \sin\theta \cos\theta \int_0^{2\pi} d\varphi I(x, 0, \theta, \varphi) \right. \\ &\quad + \int_{\cos\theta > 0} d\theta \sin\theta \cos\theta \int_0^{2\pi} d\varphi I(x, Z, \theta, \varphi) \\ &\quad \left. + \int_{\cos\theta < 0} d\theta \sin\theta |\cos\theta| \int_0^{2\pi} d\varphi I(x, y, 0, \theta, \varphi) \right] \\ &= \int_{-X_L}^{X_R} dx \{-F_0 \cos\Theta + j_B(x) + j_T(x)\}. \end{aligned} \quad (15)$$

One finds the difference of the last term at the left side and the scattering integral

$$\begin{aligned} & \int_0^Z dz \int_{-X_L}^{X_R} dx \int_0^\pi d\theta \sin\theta \int_0^{2\pi} d\varphi \left[ \sigma_{ext} I - \frac{\sigma_{sca}}{4\pi} \int_0^\pi d\theta' \sin\theta' \int_0^{2\pi} d\varphi' I(x, z, \theta', \varphi') p(x, z, \chi) \right] \\ &= \int_0^Z dz \int_{-X_L}^{X_R} dx \int_0^\pi d\theta \sin\theta \int_0^{2\pi} d\varphi I(x, z, \theta, \varphi) \left[ \sigma_{ext} - \frac{\sigma_{sca}}{4\pi} \int_0^\pi d\theta' \sin\theta' \int_0^{2\pi} d\varphi' p(x, z, \chi) \right] \\ &= \int_0^Z dz \int_{-X_L}^{X_R} dx c(x, z) \end{aligned} \quad (16)$$

The transformation of the first line in this equality to the second one is carried out via variable changes; the transformation of the second line to the third one is based upon the normalization condition (3).

Combining the equalities (14)–(16), one obtains the required balance relation

$$\begin{aligned} & \int_0^Z dz \{j_R(z) + j_L(z)\} + \int_{-X_L}^{X_R} dx \{j_B(x) + j_T(x)\} + \int_0^Z dz \int_{-X_L}^{X_R} dx c(x, y, z) \\ &= (X_L + X_R) F_0 \cos\Theta, \end{aligned} \quad (17)$$

having obvious physical sense. The term  $(X_L + X_R) F_0 \cos\Theta$  at the right side defines the full energy of radiation, entering the section across its top boundary. The first integral at the left side is equal to the difference of energy of radiation leaving the section and entering it across vertical boundaries  $x = X_R$  and  $x = -X_L$ . The second integral is equal to energy of radiation leaving the region via the bottom boundary  $z = Z$  and the top  $z = 0$ . The double integral at the left side defines energy absorbed within the section. Eq. (17) is a major result of this work.

Let us assume now that the cross section is infinite and homogeneous over  $x$ ; such a model corresponds to a light scattering plan-parallel layer of height  $Z$ . Then intensity does not depend on  $x$  and the balance relation (17) has a form (as  $-X_L \rightarrow -\infty, X_R \rightarrow \infty$ )

$$j_B + j_T + \int_0^Z dz c(z) = F_0 \cos\Theta, \quad (18)$$

196 where

197  $F_0 \cos\Theta$  energy of radiation, entering the layer across an  
 198 unit area of its top boundary,  
 199  $j_T$  energy of radiation, reflected by an unit area of the  
 200 top boundary,  
 201  $j_B$  energy of radiation, transmitted by an unit area of  
 202 the bottom boundary,

203  $\int_0^z dz c(z)$  energy of radiation absorbed in the layer.

204 Let us introduce the reflectance and transmittance for the  
 205 slab layer

$$R = \frac{j_T}{F_0 \cos\Theta}, \quad T = \frac{j_B}{F_0 \cos\Theta}. \quad (19)$$

208 Then the radiative balance relation (18) can be written as  
 209

$$R + T + A = 1, \quad (20)$$

210 where

$$A = k \int_0^z dz \sigma_{abs}(z) \Phi(z) \quad (21)$$

213 is the absorptance and  $k = 1/F_0 \cos\Theta$ . Usually, the absorptance  
 214 is found as a difference  $1 - R - T$  from calculated values of the  
 215 reflectance and transmittance. However, it can result in errors  
 216 as  $A \rightarrow 0$ . Therefore, calculations using Eq. (21) are more  
 217 accurate. It follows for a vertically homogeneous layer from  
 218 Eq. (21):

$$A = k \langle \Phi \rangle \tau_{abs} \quad (22)$$

220 or alternatively

$$A = k \Phi(z_0) \tau_{abs}, \quad (23)$$

**Table 1**

Sizes of boundary layers  $x_r^{aer}$  and  $x_r^{cloud}$  (km), where 3D effects exist, for reflectance  $R(x)$  at various values of the single scattering albedo  $\omega_0^{cloud}$  and the solar azimuth  $\Phi$ . The solar zenith angle is equal to  $60^\circ$ .

$\omega_0^{cloud}$	$\Phi = 180^\circ$		$\Phi = 0^\circ$	
	$x_r^{aer}$	$x_r^{cloud}$	$x_r^{aer}$	$x_r^{cloud}$
0.9	20.25	0.90	6.25	5.25
0.95	18.75	1.70	9.25	6.25
0.99	16.75	3.30	13.25	7.25
1.0	15.25	3.90	14.25	7.75

where

$$\tau_{abs} = \sigma_{abs} Z, \quad \langle \Phi \rangle = \frac{1}{Z} \int_0^Z \Phi(z) dz \approx \Phi(z_0) \quad (24)$$

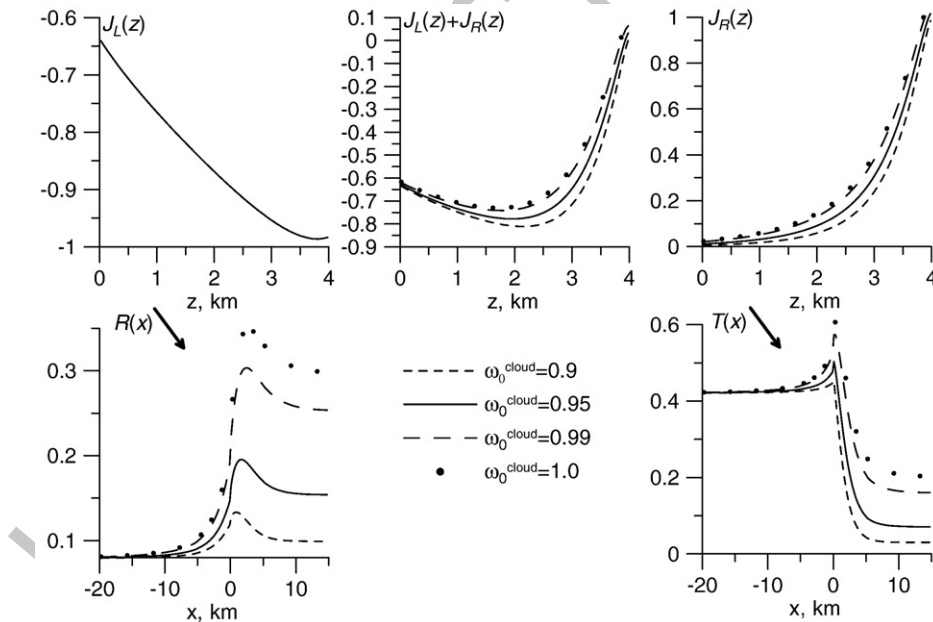
and we used the theorem on average. It follows from these equations that the average actinic flux in the scattering layer can be calculated as

$$\langle \Phi \rangle = \frac{1 - R - T}{k \tau_{abs}}. \quad (25)$$

The actinic flux is a very important parameter for atmospheric chemistry studies because it determines the photolysis rate, which is influenced by 3D effects as well. The same applies to the heating rate and the spectral density of radiation  $u = \Phi/c_0$ , where  $c_0$  is the speed of light.

#### 4. Numerical results

In this section we present numerical results for the spatial distribution of radiation absorbed, reflected and transmitted



**Fig. 2.** The normalized horizontal fluxes  $J_R(z)$ ,  $J_L(z)$ , their sum  $J_L(z) + J_R(z)$ , transmittance  $T(x)$  and reflectance  $R(x)$  for illumination from the clear sky side ( $\Phi = 0^\circ$ ).



**Table 2**

Sizes of boundary layers  $x_r^{\text{aer}}$  and  $x_r^{\text{cloud}}$  (km), where 3D effects exist, for transmittance  $T(x)$  for various values of the single scattering albedo  $\omega_0^{\text{cloud}}$  and the solar azimuth  $\Phi$ . The solar zenith angle is equal to  $60^\circ$ .

$\omega_0^{\text{cloud}}$	$\Phi = 180^\circ$		$\Phi = 0^\circ$	
	$x_r^{\text{aer}}$	$x_r^{\text{cloud}}$	$x_r^{\text{aer}}$	$x_r^{\text{cloud}}$
0.9	8.75	0.10	0.50	7.75
0.95	8.25	2.30	2.30	8.25
0.99	7.25	4.10	4.50	8.75
1.0	6.75	4.70	4.90	8.75

by a clear sky–cloud system. The following parameters are assumed in calculations:

- Zenith angle of the sun  $\Theta$ :  $60^\circ$ .
- Azimuth  $\Phi$ :  $0^\circ$ ,  $180^\circ$ .
- Height of the system:  $Z = 4$  km.
- Width of the system:  $X_L = 60$  km,  $X_R = 30$  km.
- Extinction coefficient of a clear sky region (no absorption assumed):  $\sigma_{\text{ext}}^{\text{aer}} = 0.125 \text{ km}^{-1}$ .
- Extinction coefficient in a cloud  $\sigma_{\text{ext}}^{\text{aer}} = 2.5 \text{ km}^{-1}$ .
- Single scattering albedo in a cloud  $\omega_0^{\text{cloud}}$ : 0.9, 0.95, 0.99, 1.0.

Both cloud and aerosol are assumed to be horizontally homogeneous media. Aerosol scattering is simulated using the Henyey–Greenstein phase function with the asymmetry parameter  $g=0.7$ , being decomposed into  $P_{23}$  Legendre series. Cloud C1 phase function is obtained via Mie theory for the wave-length 412 nm (van de Hulst, 1980). Its peak is replaced by delta-function using the delta-M method (Wiscombe, 1977), and the remaining regular component is presented by  $P_{27}$  Legendre series.

The solution of the problem under consideration (the light field intensity  $I(x, z, \theta, \varphi)$ ) is found via the discrete ordinate method using the code RADUGA (Nikolaeva et al., 2005) assuming the underlying black surface. This code is based upon mesh schemes of the discrete ordinate method, when regular meshes over spatial and angular variables are introduced and the transport equation in each mesh is approximated by algebraic relations. They are the balance equalities, that resulted from the integration of the radiative transfer equation over a mesh (not whole region), and the piece-linear approximation to a solution desired in a mesh. Next the successive-orders-of-scattering (SOS) iteration process with the diffusion-synthetic acceleration (DSA) method (Marchuk

and Lebedev, 1986; Larsen, 1982), where acceleration correction to a solution after each SOS iteration is found as the diffusion equation solution, are used to resolve this algebraic system. Note that the adaptive irregular spatial meshes, the ray tracing method to construct mesh approximation and the SOS iteration method with interpolation acceleration technique to resolve an approximate system are employed in the known SHDOM code (Evans, 1998).

We show the numerical results for the following functions (see Figs. 2–7):

1.  $J_R(z) = \frac{j_R(z)}{F_0 \cos \Theta}$ ,  $J_L(z) = \frac{j_L(z)}{F_0 \cos \Theta}$  – normalized horizontal fluxes,
2.  $R(x) = \frac{j_T(x)}{F_0 \cos \Theta}$ ,  $T(x) = \frac{j_B(x)}{F_0 \cos \Theta}$  – reflectance and transmittance,
3.  $C(x, z) = \frac{c(x, z)}{F_0 \cos \Theta}$  – absorptance,
4.  $\varepsilon_R(x) = 1 - R(x) / R^{\text{IPA}}(x)$ ,  $\varepsilon_T(x) = 1 - T(x) / T^{\text{IPA}}(x)$ .

where  $R^{\text{IPA}}(x)$  and  $T^{\text{IPA}}(x)$  are obtained via the IPA method.

The illumination directions are depicted in all figures by arrows. The illumination from a clear sky side corresponds to the azimuth  $\Phi = 0^\circ$ . The azimuth  $180^\circ$  is for the illumination from a cloud side.

Functions  $\varepsilon_R(x)$  and  $\varepsilon_T(x)$  are not close to zero only in the vicinity of the boundary separating clear sky and a cloud. They have jumps in vicinity of the media boundary by definition (26) due to the fact that IPA coefficients  $R^{\text{IPA}}$  and  $T^{\text{IPA}}$  are very different for a cloud (for  $x=0^+$ ) and a clear sky (for  $x=0^-$ ), whereas 2D coefficients  $R(x)$  and  $T(x)$  are smooth functions.

The dimensions of regions, where 3D effects are of importance ( $x_r^{\text{aer}}$ ,  $x_r^{\text{aer}}$ ,  $x_r^{\text{cloud}}$ , and  $x_r^{\text{cloud}}$ ) can be found from the following equations

$$\max_{x < -x_r^{\text{aer}}} |\varepsilon_R(x)| = 0.05, \quad \max_{x < -x_r^{\text{aer}}} |\varepsilon_T(x)| = 0.05 \text{ – in an aerosol} \quad (27a)$$

$$\max_{x > x_r^{\text{cloud}}} |\varepsilon_R(x)| = 0.05, \quad \max_{x > x_r^{\text{cloud}}} |\varepsilon_T(x)| = 0.05 \text{ – in a cloud.} \quad (27b)$$

They are presented in Tables 1 and 2.

In the case  $\Theta = 60^\circ$ ,  $\Phi = 0^\circ$ , and  $\omega_0^{\text{cloud}} = 1$  (the illumination from a clear sky side, see Figs. 2 and 3), the flux  $J_R(z)$  is

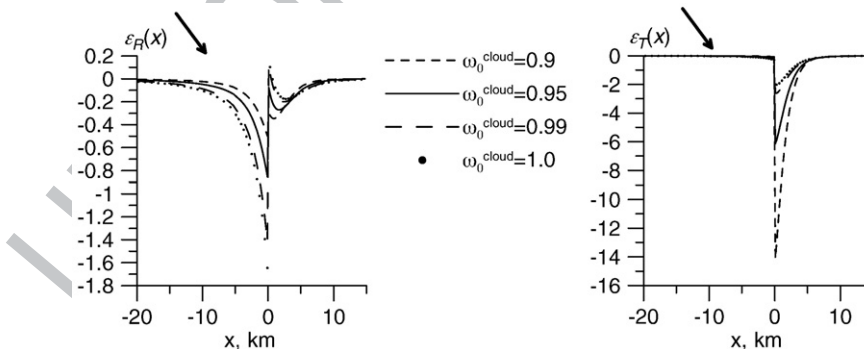
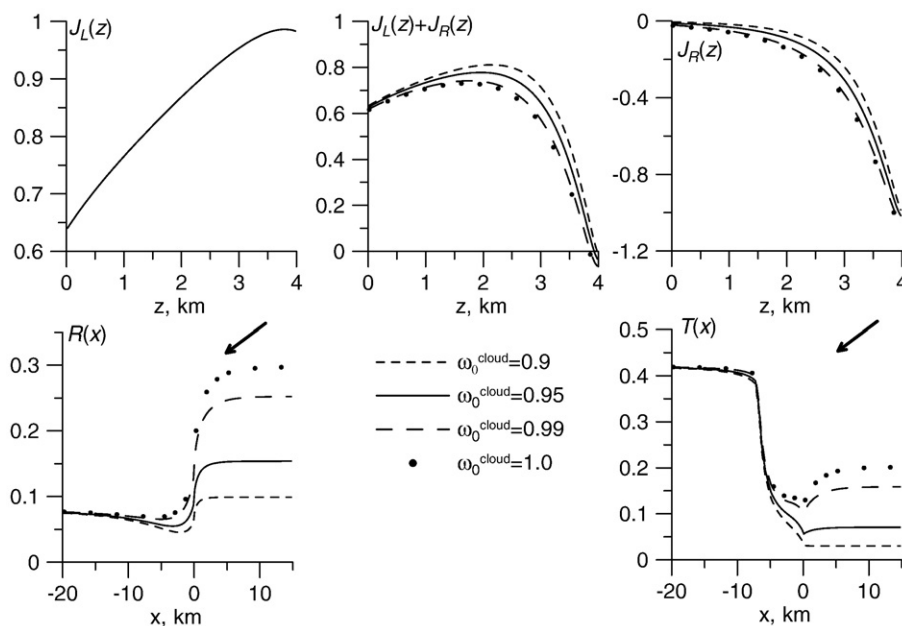


Fig. 3. Functions  $\varepsilon_R(x)$  and  $\varepsilon_T(x)$  for illumination from the clear sky side ( $\Phi = 0^\circ$ ).



**Fig. 4.** The normalized horizontal fluxes  $J_R(z)$  and  $J_L(z)$ , their sum  $J_L(z) + J_R(z)$ , transmittance  $T(x)$  and reflectance  $R(x)$  for illumination from the cloudy portion of the sky ( $\phi = 180^\circ$ ).

positive, the flux  $J_L(z)$  is negative, and the sum of these fluxes  $J_L(z) + J_R(z)$  is negative almost everywhere. So the most radiation moves from a clear sky to a cloud (in the direction of the direct light) via the vertical boundaries of the cross section, but more radiation enters the region than leaves it. “Superfluous” radiation leaves the cross section near the vertical boundary forming maxima in transmittance  $T(x)$  and reflectance  $R(x)$  (brightening effect).

Let us consider now the case at  $\phi = 180^\circ$  (illumination from a cloud side, see Figs. 4 and 5). The flux  $J_R(z)$  is negative, the flux  $J_L(z)$  is positive, the sum of these fluxes  $J_L(z) + J_R(z)$  is positive almost everywhere. So the most photons move from a cloud to a clear sky region (in the direction of direct light) via the vertical boundaries of the cross section, but more radiation leaves the region rather than enters it. This produces minima of transmittance  $T(x)$  and reflectance  $R(x)$  near the vertical boundary (shadowing effect).

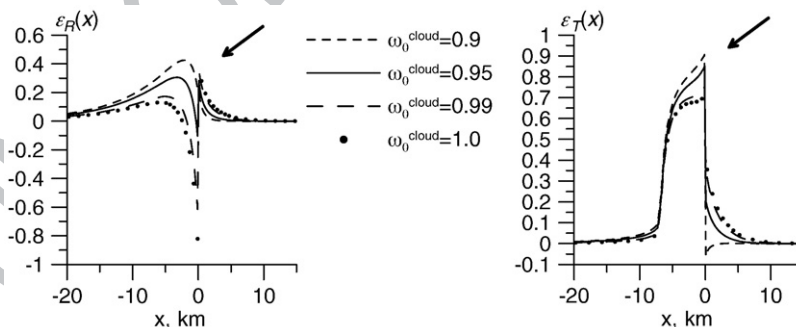
Note, if both zones were identical, sum  $J_L(z) + J_R(z)$  should be equal to zero and transmittance  $T(x)$  and reflectance

$R(x)$  should be constant functions of  $x$  (no brightening and shadowing effects).

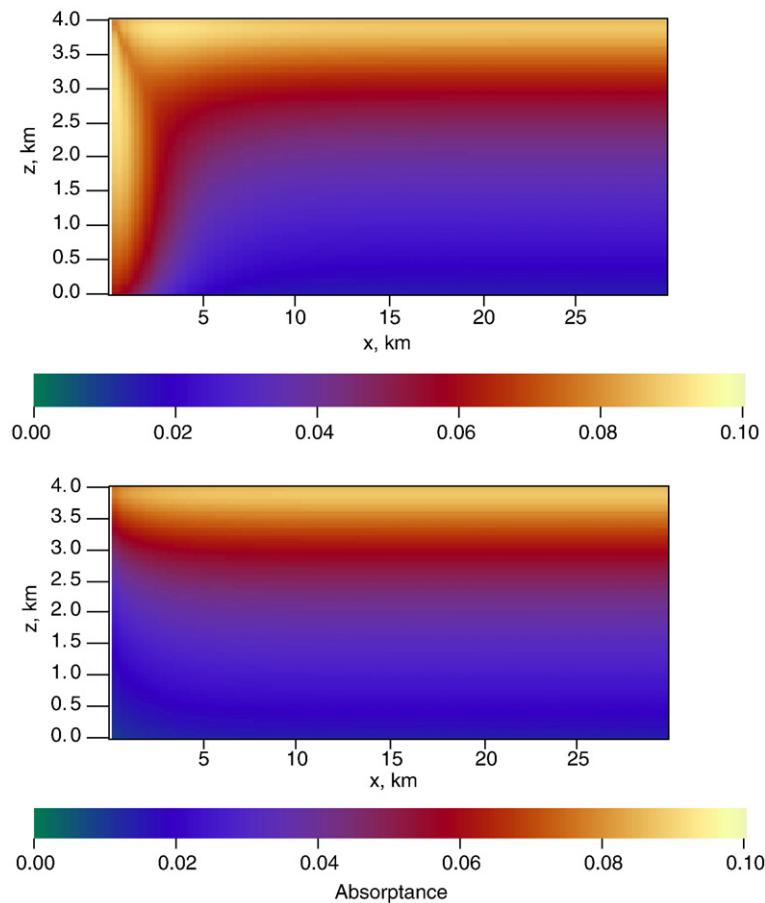
If single scattering albedo  $\omega_0^{\text{cloud}} < 1$ , some radiation is absorbed in a cloud. This process leads to the decrease of the reflectance and transmittance in the aerosol boundary layer, but does not change these parameters of light field far away from the boundary of the medium, as one might expect. We found that the decrease of the cloud single scattering albedo leads to the reduction of the brightening effect. The shadowing effect is enhanced then. This leads to the increase/decrease of the sizes of boundary layers defined above (in a clear sky) if  $\omega_0^{\text{cloud}}$  decreases (see Tables 1 and 2).

The decrease of single scattering albedo  $\omega_0^{\text{cloud}}$  is the reason why the reflectance  $R(x)$  and transmittance  $T(x)$  decrease in all points at the top and bottom boundaries of a cloud, correspondingly (far away from the interior boundary and in its vicinity).

The absorptance  $C(x, z)$  for two cases of illumination ( $\phi = 0^\circ$  and  $\phi = 180^\circ$ ) is presented in Fig. 6. It can be seen



**Fig. 5.** Functions  $\epsilon_R(x)$  and  $\epsilon_T(x)$  for illumination from the cloudy portion of the sky ( $\phi = 180^\circ$ ).



**Fig. 6.** Absorptance  $C(x,z)$  in a cloud at  $\omega_0^{\text{cloud}} = 0.99$  and illumination from the clear sky side (top) and from the cloud side (bottom). The solar zenith angle is equal to  $60^\circ$  in both cases.

the greatest absorption, and, therefore, the actinic flux and the photolysis rate, takes place near planes, across which the direct light from the sun enters a cloud. They are the top boundary (for both cases) and the left boundary (for illumination from a clear sky side). The smallest absorption takes place near the bottom cloud boundary, where the light intensity is the smallest. It follows that  $C(x,z)$  does not depend on  $x$  in deep layers of a cloud far away from the boundaries. We conclude that broken cloud fields significantly modify the radiative fluxes, the actinic flux and the photolysis rates at the regions close to the cloud boundaries.

## 5. Conclusion

The technique to study the radiation intensity distribution in the horizontally inhomogeneous clear sky–cloud system is presented. The technique is applied to a model problem, in which each medium (both an aerosol and a cloud) is homogeneous and semi-infinite along the horizontal coordinates and media have one common vertical plane boundary. All coefficients are calculated using the spatial-angular distribution of radiation intensity obtained by the code RADUGA (Nikolaeva et al., 2005) in the framework of the 2D  $(x, z)$  geometry. It is shown how the horizontal “jump” of optical properties leads to the well-known brightening and shadow-

ing effects at the edge of a cloud and in a neighboring clear sky region.

It is shown that these effects are tightly connected with the horizontal radiation transport in broken cloud systems and appear even for nonabsorbing media. The absorption of radiation can enhance or relax these effects in dependence, in particular, on the solar zenith angle.

It should be stressed that we consider only the model problem with the sharp plane vertical interior boundary and not general broken cloud systems. The optical properties of real atmosphere are often changed smoothly and there are intermediate sub-regions with varying concentrations of drops and wet aerosol particles (Redemann et al., 2009). Therefore, the results presented here are only approximately valid for the real-world cloud systems. However, we do hope that they will be found to be useful for better understanding the multi-dimensional radiative transfer in cloudy atmospheres.

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